Climate change: recent changes in sea level and the ocean



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I. Is climate change happening?2. Why care about the ocean?3. What might happen in the future?



#### Implied radiative heating of 1.67 Wm<sup>-2</sup> range: 0.6 to 2.6 Wm<sup>-2</sup>

How many light bulbs would be lit in this room by this energy input?

If room is  $20m \times 30m = 600 m^2$ 

warming of 1.67 Wm<sup>-2</sup> equivalent to

10 bulbs



#### Extreme events: Heat wave in summer 2003



**FAQ 9.1, Figure 1.** Summer temperatures in Switzerland from 1864 to 2003 are, on average, about 17°C, as shown by the green curve. During the extremely hot summer of 2003, average temperatures exceeded 22°C, as indicated by the red bar (a vertical line is shown for each year in the 137-year record). The fitted Gaussian distribution is indicated in green. The years 1909, 1947 and 2003 are labelled because they represent extreme years in the record. The values in the lower left corner indicate the standard deviation (G) and the 2003 anomaly normalised by the 1864 to 2000 standard deviation (T'/G). From Schär et al. (2004).

Summer

 $\sigma = 0.94^{\circ}C$ T'/ $\sigma = 5.4$ 

12

Frequency

10

# Modeled variability with & without anthropogenic forcing





Figure 9.5. Comparison between global mean surface temperature anomalies (°C) from observations (black) and AOGCM simulations forced with (a) both anthropogenic and natural forcings and (b) natural forcings only. All data are shown as global mean temperature anomalies relative to the period 1901 to 1950, as observed (black, Hadley Centre/Climatic Research Unit gridded surface temperature data set (HadCRUT3); Brohan et al., 2006) and, in (a) as obtained from 58 simulations produced by 14 models with both anthropogenic and natural forcings. The multi-

#### Reduction in snow cover and Arctic sea ice in summer





Observed sea ice September 2003

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## 2. Why care about the ocean?

- upper 2.5 m of ocean holds as much heat as overlying atmosphere
- 80% of anthropogenic extra heating has gone into the ocean

Oceans have been absorbing more than 80% of heat added to the climate system.



Figure 5.1. Time series of global annual ocean heat content (10<sup>22</sup> J) of 3650 to 7960h lay 8.96% black 20ve is 976ated 1080 evitu 986at. (20098), with 3955had 2000 pre-2005 senting the 90% confidence interval. The red and green curves are updates of the analyses by Ishii et al. (2006) and Willis et 9142004, over 0 to 750 m) respectively, with the error bars denoting the 90% confidence interval. The black and red curves denote the deviation from the 1961 to 1990 average and the shorter green curve denotes the deviation from the average of the black curve for the period 1993 to 2003.



### Ocean heat content change



Figure 1. Change in ocean heat content (10<sup>20</sup>J; red represents a gain in heat for the later period) between the twenty-year periods 1980-2000 and 1950-1970 diagnosed from (a) historical data integrated over the water column and (b) 1.4° ocean model output using realistic ECMWF wind



#### Ocean uptake of carbon

• 50 - 70 times as much carbon in the oceans as in the atmosphere

air-sea flux of CO<sub>2</sub> into the ocean (Takahashi et al., 2002)

Pattern reflects combination of physics & biology

annual-mean  $CO_2$  flux (mol m<sup>-2</sup>y<sup>-1</sup>)





• 1/2 the recent industrial emissions of carbon has gone into ocean



**Figure 5.9.** Changes in surface oceanic  $pCO_2$  (left; in  $\mu$ atm) and pH (right) from three time series stations: Blue: European Station for Time-series in the Ocean (ESTOC, 29°N, 15°W; Gonzalez-Dávila et al., 2003); green: Hawaii Ocean Time-Series (HOT, 23°N, 158°W; Dore et al., 2003); red: Bermuda Atlantic Time-series Study (BATS, 31/32°N, 64°W; Bates et al., 2002; Gruber et al., 2002). Values of pCO<sub>2</sub> and pH were calculated from DIC and alkalinity at HOT and BATS; pH was directly measured at ESTOC and  $pCO_2$  was calculated from pH and alkalinity. The mean seasonal cycle was removed from all data. The thick black line is smoothed and does not contain variability less than 0.5 years period.



## 3. What might happen in the future?



Figure SPM.5. Solid lines are multi-model global averages of surface warming (relative to 1980–1999) for the scenarios A2, A1B and B1, shown as continuations of the 20th century simulations. Shading denotes the ±1 standard deviation range of individual model annual averages. The orange line is for the experiment where concentrations were held constant at year 2000 values. The grey bars at right indicate the best estimate (solid line within each bar) and the **likely** range assessed for the bix SRES marker scenarios. The assessment of the best estimate and **likely** ranges in the grey bars includes the AOGCMs in the left part of the figure, as well as results from a hierarchy of independent models and observational constraints. [Figures 10.4 and 10.29]

#### What happens if we burn all our fossil fuels?



- Initial fast rise in atmospheric CO<sub>2</sub>
- Ocean (& terrestrial) uptake
- Eventually approach a steady state

steady state set by the total amount of carbon emitted



$$\Delta F_{CO2} \approx \frac{\alpha}{I_B} \Delta C$$

PhD thesis of Phil. Goodwin (Liverpool) If burn all of the conventional fossil fuels, 5000GtC, then extra heating of 8.5 Wm<sup>-2</sup>

Same amount of heating given in this room (20m x 30m) by these 100 W light bulbs:

50 bulbs

## Summary

Evidence for warming of the climate system is unequivocal.

We are seeing coherent changes in many aspects of the climate system other than temperature.

The role of greenhouse gases is well understood and their increases are clearly identified.

The net effect of human activities is now quantified and known to cause a warming at the Earth's surface.



#### **RADIATIVE FORCING COMPONENTS**

**Figure SPM.2.** Global average radiative forcing (RF) estimates and ranges in 2005 for anthropogenic carbon dioxide ( $CO_2$ ), methane ( $CH_4$ ), nitrous oxide ( $N_2O$ ) and other important agents and mechanisms, together with the typical geographical extent (spatial scale) of the forcing and the assessed level of scientific understanding (LOSU). The net anthropogenic radiative forcing and its range are also shown. These require summing asymmetric uncertainty estimates from the component terms, and cannot be obtained by simple addition. Additional forcing factors not included here are considered to have a very low LOSU. Volcanic aerosols contribute an additional natural forcing but are not included in this figure due to their episodic nature. The range for linear contrails does not include other possible effects of aviation on cloudiness. {2.9, Figure 2.20}

